

Uncertainties and limitations of using carbon-13 and oxygen-18 leaf isotope exchange to estimate the temperature response of mesophyll CO₂ conductance in C₃ plants

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Summary

- The internal CO₂ gradient imposed by mesophyll conductance (g_m) reduces substrate availability for C₃ photosynthesis. With several assumptions, estimates of g_m can be made from coupled leaf gas exchange with isoflux analysis of carbon $\Delta^{13}\text{C}-g_m$ and oxygen in CO₂, coupled with transpired water (H₂O) $\Delta^{18}\text{O}-g_m$ to partition g_m into its biochemical and anatomical components. However, these assumptions require validation under changing leaf temperatures.
- To test these assumptions, we measured and modeled the temperature response (15–40°C) of $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ along with leaf biochemistry in the C₃ grass *Panicum bisulcatum*, which has naturally low carbonic anhydrase activity.
- Our study suggests that assumptions regarding the extent of isotopic equilibrium (θ) between CO₂ and H₂O at the site of exchange, and that the isotopic composition of the H₂O at the sites of evaporation ($\delta_{w-e}^{18}\text{O}$) and at the site of exchange ($\delta_{w-ce}^{18}\text{O}$) are similar, may lead to errors in estimating the $\Delta^{18}\text{O}-g_m$ temperature response. The input parameters for $\Delta^{13}\text{C}-g_m$ appear to be less sensitive to temperature. However, this needs to be tested in species with diverse carbonic anhydrase activity.
- Additional information on the temperature dependency of cytosolic and chloroplastic pH may clarify uncertainties used for $\Delta^{18}\text{O}-g_m$ under changing leaf temperatures.

Introduction

Diffusional limitations to CO₂ movement into and within a leaf result in reduced CO₂ availability at the site of carboxylation and can therefore limit rates of photosynthesis (Evans & von Caemmerer, 1996). The initial resistance to CO₂ diffusion through stomata from the leaf surface (C_a) to the intercellular air spaces (C_i) is well characterized and is known to strongly influence rates of photosynthesis (Warren, 2008). Within a leaf, CO₂ must further diffuse from the intercellular air spaces to the site of carboxylation inside chloroplast for CO₂ fixation by the enzyme Rubisco. These final steps of the CO₂ diffusion pathway are generally referred to as mesophyll CO₂ conductance g_m (Evans & von Caemmerer, 1996).

In C₃ plants, g_m has been shown to vary between species (von Caemmerer & Evans, 2015), to acclimate under different environmental growth conditions (Flexas *et al.*, 2008), and to change with leaf age (Niinemets *et al.*, 2006; Barbour *et al.*, 2016). Additionally, g_m in C₃ plants has been shown to vary in response to short-term changes in measurement temperatures, light, and CO₂ concentration in some but not all species (Flexas *et al.*, 2007; Douthe *et al.*, 2011; Tazoe *et al.*, 2011; Evans & von

Caemmerer, 2013; von Caemmerer & Evans, 2015). Leaf properties, such as the arrangement and compactness of mesophyll cells, chloroplast orientation to the intercellular airspaces, and cell wall and membrane properties, have all been proposed to influence adaptive and long-term acclimation of g_m , whereas differences in leaf biochemistry (e.g. carbonic anhydrase (CA) and aquaporins) are thought to potentially drive dynamic g_m responses to short-term changing environments (Gillon & Yakir, 2000; Evans *et al.*, 2009; von Caemmerer & Evans, 2015). However, our understanding of how leaf anatomy and biochemistry influence g_m is incomplete, primarily because there are no direct ways to measure g_m and the contribution of these various components.

Historically, in C₃ plants, combined measurement of g_m using isoflux of carbon (C) in CO₂ ($\Delta^{13}\text{C}-g_m$) and isoflux of oxygen (O) in CO₂ and transpired water (H₂O) ($\Delta^{18}\text{O}-g_m$) have been used to partition mesophyll conductance into wall conductance g_w (i.e. cell wall, plasma membrane, and cytosol) and chloroplast conductance g_{ch} (i.e. chloroplast membrane and stroma) (Evans *et al.*, 1994; Gillon & Yakir, 2000). Unfortunately, there are several assumptions needed to derive g_m from both the $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ methods, and only a few studies have simultaneously

combined both methods on a limited number of species (Gillon & Yakir, 2000; Barbour *et al.*, 2016) under a few short-term environmental conditions (irradiance and humidity in cotton; Loucos *et al.*, 2017). Therefore, a critical comparison of the assumptions used to calculate $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$, particularly in response to temperature, is needed.

For $\Delta^{13}\text{C}-g_m$, the Rubisco fractionation factor b is a key variable (Evans *et al.*, 1986). Estimation of b is difficult; consequently, there are only a few reports on b , and uncertainties remain in its temperature dependency (O'Leary *et al.*, 1992; Tcherkez & Farquhar, 2005; Evans & von Caemmerer, 2013). In addition, in the absence of a species-specific temperature response of the CO_2 compensation point at the site of carboxylation (Γ^*) and gas-exchange measurements at 2% $[\text{O}_2]$, there could be potential uncertainties for the temperature dependency of fractionation factors associated with respiration and photorespiration, respectively (Evans & von Caemmerer, 2013).

Alternatively, for the $\Delta^{18}\text{O}-g_m$ calculations, it is unclear if the CO_2 and the H_2O at the site of exchange are in full isotopic equilibrium θ and if the isotopic signature of the H_2O at the site of evaporation δ_{w-e}^{18} accurately represents the H_2O signature at the site of exchange δ_{w-ce}^{18} . The assumption that there is a full isotopic equilibrium between CO_2 and H_2O at the site of exchange (e.g. $\theta = 1$) is primarily estimated by the activity of leaf CA. However, published CA activity varies widely between studies, species, tissue collection methods, and growth conditions (Hatch & Burnell, 1990; Gillon & Yakir, 2000; Cousins *et al.*, 2008). Boyd *et al.* (2015) suggested deactivation of CA activity in *Setaria viridis* at temperatures above 25°C. In addition, the influence of temperature-induced changes in pH on CA activity cannot be ruled out. Therefore, changes in leaf temperature may offset θ from 1, and this may be higher in species with low CA activity, particularly above 25°C. The assumption that the $\delta_{w-ce}^{18} = \delta_{w-e}^{18}$ has been justified because the distance between the outer cell wall and the chloroplast appressed to the intercellular airspace is short and may lead to the only small gradient in H_2^{18}O enrichment (Gillon & Yakir, 2000; Barbour *et al.*, 2016). However, changing leaf temperatures may change H_2O flux inside the leaf and potentially the location that H_2O transitions between the liquid and vapor phase (Buckley *et al.*, 2017). Taken together, assumptions regarding parameter values in calculations of $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ may propagate uncertainties in estimating g_m in C_3 plants as leaf temperature changes.

Assuming that the calculations of $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ are parameterized correctly, it has been suggested that $\Delta^{13}\text{C}-g_m$ provides estimates of the total mesophyll conductance from the intercellular air spaces to the chloroplast stroma, whereas $\Delta^{18}\text{O}-g_m$ estimates internal CO_2 conductance to the chloroplast surface (Gillon & Yakir, 2000; Barbour *et al.*, 2016). Accordingly, the $\Delta^{13}\text{C}-g_m$ is often expected to be $c. 0.66 \times \Delta^{18}\text{O}-g_m$ (Yakir, 1998). However, short-term changes in leaf temperatures affect the rate of diffusional processes and biochemical reactions; hence, temperature affects CO_2 diffusion through membranes and liquid path, H_2O fluxes, and CO_2 - H_2O equilibrium within a leaf (Evans & von Caemmerer, 2013; Barbour *et al.*, 2016). Therefore, investigating temperature dependency of $\Delta^{18}\text{O}-g_m$ coupled

with the $\Delta^{13}\text{C}-g_m$ provides an opportunity to test the assumptions associated with estimating $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$.

Here, we simultaneously determined the temperature response of $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ to test the assumptions used for $\Delta^{13}\text{C}-g_m$ against $\Delta^{18}\text{O}-g_m$, and vice versa. For this, we measured photosynthetic $\Delta^{13}\text{C}$ and $\Delta^{18}\text{O}$ under changing leaf temperatures, leaf CA activities, and the pH response of CA activity. We used the C_3 grass *Panicum bisulcatum*, which has high rates of CO_2 assimilation and stomatal conductance but naturally low CA activity. Comparison of the temperature responses of $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ suggests that errors in assumptions of parameters in $\Delta^{18}\text{O}-g_m$ calculations will strongly influence g_m estimates. By contrast, calculations of $\Delta^{13}\text{C}-g_m$ appeared to be more temperature robust.

Materials and Methods

Plant material and growth environment

Panicum bisulcatum (PI286485) seeds were germinated in a commercial Sun Gro® Sunshine® LC1 Grower Mix with RESILIENCE™ (<http://www.bfgsupply.com>) at the Washington State University, Pullman, WA, USA, in a controlled-environment growth cabinet (model GC-16; Enconair Ecological Chambers Inc., Winnipeg, MB, Canada). Growth conditions were set at 16 h photoperiod including a 2 h ramp at the beginning and at the end of the light period and maximum photosynthetic photon flux density of $600 \mu\text{mol m}^{-2} \text{s}^{-1}$. Light and dark temperatures were maintained at 28 ± 1 and 18 ± 1 °C, respectively, and the mean relative humidity was $60 \pm 7\%$. At 2–3 wk after germination, two healthy seedlings were transplanted into a 2 l pre-irrigated pot containing grower mix used for the germination. A week later, one seedling was removed, leaving one healthy plant per pot. Subsequently, plants were watered daily to field capacity for the remainder of the experiment and received 21-5-20 fertilizer (JR Peters Inc., Allentown, PA, USA; <http://www.jrpeters.com>) with Scott-Peters Soluble Trace Element Mix (The Scotts Co., Marysville, OH, USA) twice a week at concentrations in H_2O of 2.5 g l^{-1} and 10.0 mg l^{-1} , respectively. Plant location within the growth chamber was randomized daily.

Coupled leaf gas exchange and isoflux measurements

The LI-6400XT infrared gas analyzer (Li-Cor, Lincoln, NE, USA; operating as an open system), cavity-ring down absorption spectroscope (L2130-i; Picarro Inc., Sunnyvale, CA, USA), and the tunable-diode laser absorption spectroscope (TDLAS, model TGA 220A; Campbell Scientific Inc., Logan, UT, USA) were coupled as described by Ubierna *et al.* (2017). The entire LI-6400XT, the 2 cm × 6 cm leaf chamber (6400-11, Li-Cor), and LI-6400-18-RGB light source were placed in a growth cabinet (model EF7, Conviron; Controlled Environments Inc., MN, USA). The inlet gas line to the LI-6400XT cuvette was split using a brass Tee (Swagelok® Tube Fitting, <http://www.swagelok.com>) and part of the flow was diverted to the TDLAS (reference gas). The flow from the matching tube of the LI-6400XT leaf sample cuvette (sample gas) was split between the L2130-i and the

TDLAS. Air supplied to the TDLAS was passed through a Nafion[®] dryer (PD[™]-200T-12; Perma Pure LLC, Toms River, NJ, USA), and the sample line tube (type 1300 Synflex[®]) for the L2130-i was wrapped with an electrical heating cable to avoid condensation (Kolbe & Cousins, 2018).

The TDLAS data were calibrated using the concentration series method (Tazoe *et al.*, 2011, Supporting Information 1; Ubierna *et al.*, 2013), and the TDLAS levels of precision (standard deviation) for CO₂ (¹²CO₂ + ¹³CO₂) molar fraction, δ¹³C, and δ¹⁸O were ± 0.06 μmol CO₂ mol⁻¹ dry air, ± 0.26‰, and ± 0.21‰, respectively. It should be noted that the δ¹⁸O signature was expressed to a common scale (Vienna standard mean ocean water) for comparison of absolute values with the L2130-i.

The L2130-i was calibrated using three running standards calibrated against standard light Antarctic precipitation and Puerto Rico (US Geological Survey) standards. To correct for the concentration dependency of the L2130-i measurements, a standard curve of H₂O vapor concentration ([H₂O]; ppm) of a known and constant δ_w¹⁸ was determined (Supporting Information Fig. S1a). Corrections were made for all measurements at or above 20 000 ppm (Fig. S1b). The L2130-i precision (standard deviation) for δ_w¹⁸ of H₂O vapor was ± 0.44‰.

Δ¹³C and Δ¹⁸O (the leaf C and O isotope net discrimination of CO₂, respectively) and δ_{w-out}¹⁸ (the δ¹⁸O of H₂O vapor leaving the leaf chamber) were measured at leaf temperatures of 15, 20, 25, 30, 35, and 40°C on the youngest fully expanded leaf from four plants placed in the LI-6400XT leaf chamber. All measurements started at either 25 or 30°C, and the subsequent measurement temperatures were randomly selected and controlled using both the growth cabinet and LI-6400XT temperature control systems. Throughout the measurements, the desiccant and the soda lime column of the LI-6400XT were fully bypassed. The inlet gas to the LI-6400XT was CO₂ and H₂O free, and no supplemental H₂O vapor was added to the reference gas. Therefore, the concentration and isotopic composition of H₂O vapor leaving the chamber was only determined by leaf transpiration. The leaf chamber was maintained at a CO₂ partial pressure of C_a ≈ 35 Pa, 2% [O₂] and a photosynthetic photon flux density of 1200 μmol m⁻² s⁻¹. To avoid errors associated with [O₂] in the gas-exchange calculations, the LI-6400XT program was edited for 2% O₂. Every day, before the leaf measurements, a leak test was determined on an empty LI-6400XT chamber at C_a ≈ 20 Pa. At each measurement temperature the leaves were acclimated for minimum 30 min or until stable values of A_{net} and g_s were achieved. Data were subsequently collected over the next 20 min and the LI-6400XT was set to log data only when the TDLAS analyzed the sample line.

Leaf C isotope discrimination (Δ¹³C) and leaf O isotope discrimination (Δ¹⁸O)

We use δ¹³ and δ¹⁸ for δ¹³C and δ¹⁸O values of CO₂, respectively, and the δ¹⁸O in H₂O vapor or liquid H₂O is referred to as δ_w¹⁸. The observed photosynthetic discrimination against ¹³CO₂

(Δ¹³C) and C¹⁸O¹⁶O (Δ¹⁸O) was calculated as (Evans *et al.*, 1986):

$$\Delta = \frac{\frac{C_{in}}{C_{in}-C_{out}}(\delta_{out} - \delta_{in})}{1 + \delta_{out} - \frac{C_{in}}{C_{in}-C_{out}}(\delta_{out} - \delta_{in})} \quad \text{Eqn 1}$$

(C is the ¹²CO₂ mole fraction in dry air in and out of the leaf chamber; δ refers to either the δ¹³C or the δ¹⁸O in the calculation of Δ¹³C or Δ¹⁸O, respectively).

CO₂ mesophyll conductance from Δ¹³C (Δ¹³C-g_m)

Δ¹³C-g_m was calculated from the difference between estimated C isotope discrimination for C₃ plants assuming infinite g_m (Δ_i¹³), and that measured by the LI6400XT and TDLAS coupled system (Farquhar & Cernusak, 2012):

$$\Delta_i^{13} = \frac{1}{1-t} \left(a \frac{C_a - C_s}{C_a} + a \frac{C_s - C_i}{C_a} \right) + \frac{1+t}{1-t} \left(b \frac{C_i}{C_a} - e f \frac{\alpha_b}{\alpha_{el}} \frac{R_d}{A_{net} + R_d} \frac{C_i - \Gamma^*}{C_a} - f \frac{\alpha_b \Gamma^*}{\alpha_f C_a} \right) \quad \text{Eqn 2}$$

The definition and derivation of variables are explained in Table S1. The CO₂ compensation point at the site of carboxylation, Γ* and its temperature dependency in *P. bisulcatum* was estimated according to Sharwood *et al.* (2016).

The difference between Δ_i¹³ and Δ¹³C provides mesophyll resistance r_m by (Farquhar & Cernusak, 2012):

$$r_m = \frac{1-t}{1+t} (\Delta_i^{13} - \Delta^{13}C) \frac{C_a}{A \left(b - a_m - \frac{\alpha_b}{\alpha_e} e f \frac{R_d}{A_{net} + R_d} \right)} \quad \text{Eqn 3}$$

(a_m is the fractionation during diffusion and dissolution of CO₂ through the H₂O). Note that Eqn 3 is presented in Farquhar & Cernusak (2012, Appendix 3) and used by Barbour *et al.* (2016).

It can be written as:

$$\Delta^{13}C-g_m = \frac{1}{r_m} \quad \text{Eqn 4}$$

According to Fick's law of diffusion, the leaf mesophyll [CO₂] by the ¹³C method was derived as:

$$C_{c13} = C_i - \frac{A_{net}}{\Delta^{13}C-g_m} \quad \text{Eqn 5}$$

(C_{c13} is chloroplastic [CO₂] estimated by Δ¹³C-g_m).

CO₂ mesophyll conductance from Δ¹⁸O (Δ¹⁸O-g_m)

Calculation of mesophyll conductance from Δ¹⁸O (Δ¹⁸O-g_m)
The δ¹⁸O of H₂O vapor transpired by the leaf (δ_{w-E}¹⁸) is given by

(Simonin *et al.*, 2013):

$$\delta_{w-E}^{18} = \frac{W_{out}\delta_{w-out}^{18} - W_{in}\delta_{w-in}^{18} + \frac{(\delta_{w-in}^{18} - \delta_{w-out}^{18})W_{out}W_{in}}{1000}}{W_{out} - W_{in}} \quad \text{Eqn 6}$$

where δ_{w-in}^{18} and δ_{w-out}^{18} are the $\delta^{18}\text{O}$ of H_2O vapor entering (W_{in}) and leaving (W_{out}) the leaf chamber, respectively. In the current study, dry air was used for the inlet air, so $\delta_{w-E}^{18} = \delta_{w-out}^{18}$. The $\delta^{18}\text{O}$ of liquid H_2O at the sites of evaporation within the leaf was calculated with the modified Craig–Gordon model (Bottinga and Craig, 1968):

$$\delta_{w-e}^{18} = \delta_{w-E}^{18} + \varepsilon^* + \varepsilon_k + (\delta_{w-out}^{18} - \varepsilon_k - \delta_{w-E}^{18}) \frac{e_a}{e} \quad \text{Eqn 7}$$

where ε^* is the equilibrium fractionation during H_2O evaporation from liquid to vapor and it is temperature (T_k) dependent, given as (Bottinga and Craig, 1968):

$$\varepsilon^* = 2.664 - 3.206 \left(\frac{1000}{T_k} \right) + 1.534 \left(\frac{10^6}{T_k^2} \right) \quad \text{Eqn 8}$$

The ε_k is the kinetic fractionation of H_2^{18}O diffusion from the leaf intercellular airspace to the atmosphere, which is dependent on boundary layer g_b and stomatal g_s conductance and their associated fractionation factors (Farquhar *et al.*, 1989):

$$\varepsilon_k = \frac{28g_s^{-1} + 19g_b^{-1}}{g_s^{-1} - g_b^{-1}} \quad \text{Eqn 9}$$

The $\delta^{18}\text{O}$ of CO_2 at the sites of exchange in the cytosolic δ_{ce}^{18} , assuming that H_2O at the site of exchange δ_{w-ce}^{18} is isotopically the same as H_2O at the sites of evaporation δ_{w-e}^{18} , was calculated as (Cernusak *et al.*, 2004):

$$\delta_{ce}^{18} = \delta_{w-e}^{18} \theta (1 + \varepsilon_w) + \theta \varepsilon_w + \delta_{c0} (1 - \theta) \quad \text{Eqn 10}$$

where δ_{c0} is the $\delta^{18}\text{O}$ of unreacted CO_2 , θ is isotopic equilibrium between CO_2 and H_2O , and ε_w is the temperature T_k equilibrium fractionation between chloroplast CO_2 and H_2O given as:

$$\varepsilon_w = \frac{17604}{T_k} - 17.93 \quad \text{Eqn 11}$$

The leaf mesophyll $[\text{CO}_2]$ can be calculated with the ^{18}O method (C_{m18}) solving Eqn S4 and Eqn 10 for C_{m18} and assuming complete isotopic equilibrium ($\theta = 1$), where δ_{ce}^{18} equals the $\delta^{18}\text{O}$ of cytosolic CO_2 δ_c^{18} , as (Barbour *et al.*, 2016; Ubierna *et al.*, 2017):

$$C_{m18} = \frac{C_i(\delta_i^{18} - \alpha_w^{18}\delta_A^{18} - a_w^{18})}{\delta_{ce}^{18} - \alpha_w^{18}\delta_A^{18} - a_w^{18}} \quad \text{Eqn 12}$$

where the definition and derivation of variables are explained in Methods S1 and Table S2.

According to Fick's law of diffusion:

$$\Delta^{18}\text{O-}g_m = \frac{A_{net}}{C_i - C_{m18}} \quad \text{Eqn 13}$$

Modeling temperature response of Rubisco discrimination factor b , isotopic equilibrium θ , and $\delta^{18}\text{O}$ of H_2O at the sites of evaporation (δ_{w-e}^{18})

The estimation of $\Delta^{13}\text{C-}g_m$ in Eqn 2 assumes that the Rubisco fractionation factor b is independent of the temperature. Similarly, the estimation of $\Delta^{18}\text{O-}g_m$ assumes that the local cytosolic H_2O δ_{w-ce}^{18} is isotopically similar to the H_2O at the sites of evaporation δ_{w-e}^{18} (Eqn 10) and that there is a full equilibrium between CO_2 and cytosolic H_2O ($\theta = 1$). We tested these assumptions, with the caveat that $\Delta^{18}\text{O-}g_m$ measures g_m to the chloroplast surface and that $\Delta^{13}\text{C-}g_m$ estimates g_m to the site of carboxylation, such that the inherent difference between $\Delta^{13}\text{C-}g_m$ and $\Delta^{18}\text{O-}g_m$ is primarily determined by the conductance across the chloroplast envelope and can be accounted for as $\Delta^{13}\text{C-}g_m \approx 0.66 \times \Delta^{18}\text{O-}g_m$ across temperatures. For simplicity across the leaf temperatures, we derived the optimal solution needed to minimize the residual sum of squares for b , θ , and δ_{w-e}^{18} assuming that the $\Delta^{13}\text{C-}g_m$ is equal to the $\Delta^{18}\text{O-}g_m$, and vice versa.

Leaf CA activity and pH response

Fresh leaf discs (0.71 cm²) were extracted on ice in a mortar with a pestle in 1 ml of 100 mM HEPES (pH 7.8), 1% (w/v) polyvinylpyrrolidone, 1 mM EDTA, 10 mM dithiothreitol, 0.1% (v/v) Triton X-100, and 2% (v/v) protease inhibitor cocktail (P9599; Sigma-Aldrich). Crude extracts were centrifuged at 4°C for 1 min at 17 000 g, and the supernatant was collected for immediate use in the CA assay. CA activity was measured using a membrane inlet mass spectrometer to measure the rates of $^{18}\text{O}_2$ exchange from labeled $^{13}\text{C}^{18}\text{O}_2$ to H_2^{16}O with a total C concentration of 1 mM (Silverman, 1982; Badger and Price, 1989; Hatch & Burnell, 1990). The pH response of hydration rates k_{CA} was calculated from the enhancement in the rate of ^{18}O loss over the uncatalyzed rate with the nonenzymatic first-order rate constant for the hydration of CO_2 calculated for the assay pH (6.8–8.2) at 25°C using the equation from (Jenkins, 1989). The CO_2 concentration was calculated using the temperature-appropriate pK_a assuming an ionic strength of 0.1 M (Harned & Bonner, 1945), and the $p\text{CO}_2$ was calculated using the temperature-appropriate Henry's constant (Sander, 2015). The temperature dependency of leaf CA activity CA_{leaf} was estimated at pH 8.0 using our measured temperature response of $\Delta^{13}\text{C-}g_m$ to derive chloroplast $[\text{CO}_2]$ and the temperature dependency of CA activity according to Boyd *et al.* (2015). The pH sensitivity of leaf CA was predicted using our measured pH response of k_{CA} , $\Delta^{13}\text{C-}g_m$ -derived chloroplast $[\text{CO}_2]$ and the temperature dependency of CA of Boyd *et al.* (2015).

Statistical analysis

Statistical analyses and estimation of the optimal solution for parameters were performed using R (R Core Team, 2017). The effect of temperature was compared using a linear mixed-effect model using the LME4 package (Bates *et al.*, 2015). Significance tests were performed using ANOVA ($n = 4$). Variable means were ranked using a *post hoc* Tukey test.

Results

Temperature response of gas exchange and discrimination

The rate of net CO₂ assimilation A_{net} was responsive to changes in leaf temperature from 15 to 35°C, with temperature optimum around 35°C (Fig. 1a; Table S3). Similarly, stomatal conductance g_s increased with leaf temperature from 15 to 30°C, but was unchanged above 30°C, despite increases in the leaf-to-air vapor pressure deficit at 35 and 40°C (Fig. 1b,c; Table S3). The ratio of intercellular to ambient [CO₂] C_i/C_a did not significantly

change across measurement temperatures ($P > 0.05$; Fig. 1d; Table S3); however, the transpiration rates E increased significantly ($P < 0.05$) with temperature (Fig. S2). In general, the C isotope discrimination $\Delta^{13}\text{C}$ tended to increase with temperature, except at 15°C where $\Delta^{13}\text{C}$ was similar to the 30 and 35°C values (Fig. 1e; Table S3). Conversely, the O isotope discrimination $\Delta^{18}\text{O}$ decreased with temperature, except at 40°C compared with 35°C (Fig. 1f; Table S3).

Comparison of $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$

The mesophyll conductance derived from $\Delta^{13}\text{C}$ $\Delta^{13}\text{C}-g_m$, assuming a constant Rubisco discrimination factor b of 29‰, increased significantly with leaf temperature ($P_{5,14} < 0.001$). However, the mesophyll conductance derived from $\Delta^{18}\text{O}$ $\Delta^{18}\text{O}-g_m$, assuming fully isotopic equilibrium of CO₂ with the H₂O at the site of evaporation, increased with temperature between 15 and 30°C but did not respond from 30 to 40°C (Fig. 2; Table S3). As already described, the $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ estimated did not differ between 15 and 25°C;

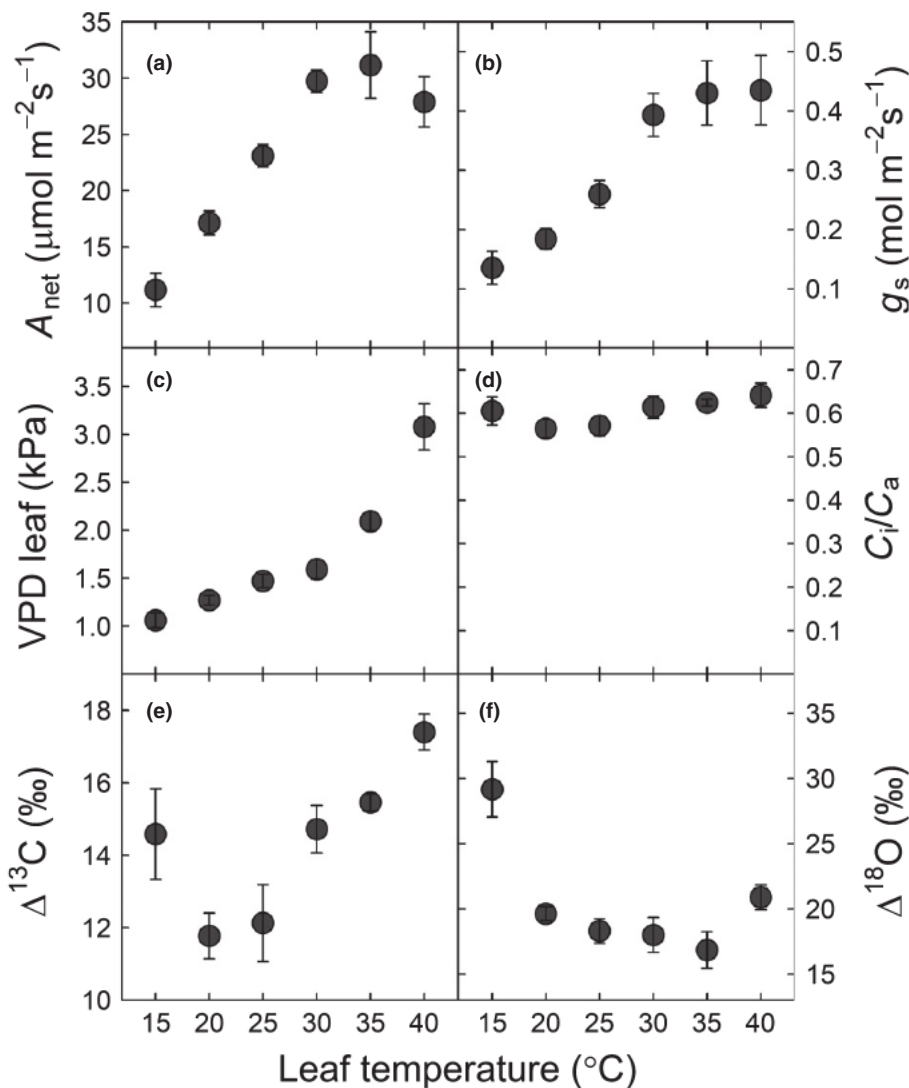


Fig. 1 The temperature response of (a) the net rate of CO₂ assimilation (A_{net}), (b) stomatal conductance to water (g_s), (c) leaf-to-air vapor pressure deficit (VPD leaf), (d) ratio of intercellular to ambient [CO₂] C_i/C_a , (e) leaf carbon isotope discrimination ($\Delta^{13}\text{C}$), and (f) leaf oxygen isotope discrimination ($\Delta^{18}\text{O}$) in *Panicum bisulcatum*. Measurements were performed at c. 35 Pa [CO₂], 1200 $\mu\text{mol m}^{-2}\text{s}^{-1}$ photosynthetic photon flux density and 2% [O₂]. Values are mean \pm SE, $n = 4$.

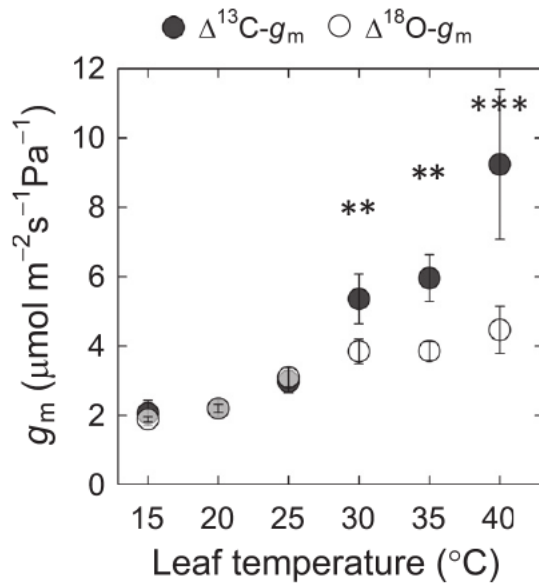


Fig. 2 The temperature response of mesophyll conductance (g_m), calculated with $\Delta^{13}\text{C}-g_m$ method assuming Rubisco fractionation factor $b = 29\text{‰}$ for all temperatures (closed circles) and $\Delta^{18}\text{O}-g_m$ method (open circles) in *Panicum bisulcatum*. For $\Delta^{18}\text{O}-g_m$ method, g_m was calculated assuming isotopic equilibrium ($\theta = 1$) and $\delta_{w-e}^{18} = \delta_{w-ce}^{18}$. Repeated measures ANOVA and pairwise comparisons between two methods across leaf temperature were used to test statistical significances: **, $P < 0.01$; ***, $P < 0.001$. Values are mean \pm SE, $n = 4$.

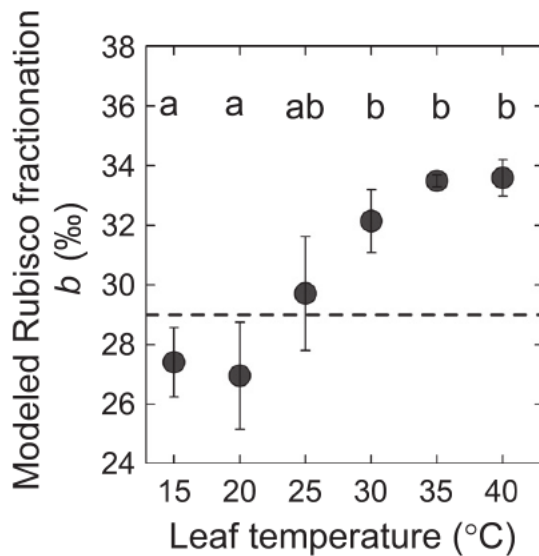


Fig. 3 Modeled changes in Rubisco fractionation factor b needed to minimize the difference between the $\Delta^{13}\text{C}-g_m$ and the measured $\Delta^{18}\text{O}-g_m$ in *Panicum bisulcatum*. Dashed line represents $b = 29\text{‰}$ used in estimating $\Delta^{13}\text{C}-g_m$ in Fig. 2. The letters are ranking (from lowest = a) for temperatures derived using a multiple-comparison Tukey *post hoc* test. Values are mean \pm SE, $n = 3$.

however, between 30 and 40°C the $\Delta^{13}\text{C}-g_m$ was significantly higher than the $\Delta^{18}\text{O}-g_m$ (Fig. 2; Table S3), suggesting uncertainty in the assumptions made for $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ across the measurement temperatures.

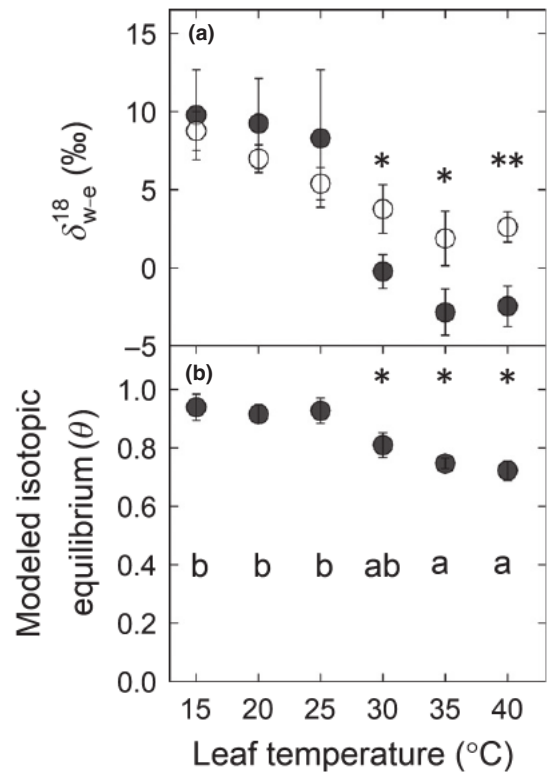


Fig. 4 The temperature response of measured $\delta^{18}\text{O}$ of the liquid water at the sites of evaporation inside the leaf δ_{w-e}^{18} (open circles) and modeled δ_{w-e}^{18} (closed circles) needed to minimize the difference between $\Delta^{18}\text{O}-g_m$ and $\Delta^{13}\text{C}-g_m$ in *Panicum bisulcatum* (a). Asterisks in (a) represent significance for pairwise comparison between two methods across leaf temperatures. The temperature response of modeled isotopic equilibrium (θ) assuming $\Delta^{18}\text{O}-g_m$ is equal to $\Delta^{13}\text{C}-g_m$ in *P. bisulcatum* (b). Asterisks in (b) represent significance for one sample *t*-test across leaf temperatures indicating measured value is statistically different from 1. The letters are ranking (from lowest = a) for temperatures derived using a multiple-comparison Tukey *post hoc* test. Values are mean \pm SE, $n = 4$. *, $P < 0.05$; **, $P < 0.01$.

Temperature response of Rubisco discrimination factor b , $\delta^{18}\text{O}$ of H_2O at the sites of evaporation δ_{w-e}^{18} and isotopic equilibrium θ

As already described, $\Delta^{13}\text{C}-g_m$ was initially estimated with a constant b across temperatures. Therefore, the b value was numerically solved for to minimize the difference between $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ at each measurement temperature, resulting in a significant change in b with temperature ($P_{5,10} < 0.01$) from 27.4 ± 1.2 to $33.6 \pm 0.6\text{‰}$ (Fig. 3). Alternatively, the $\Delta^{18}\text{O}-g_m$ presented in Fig. 2 is based on the assumption that CO_2 is fully equilibrated (i.e. $\theta = 1$) with H_2O at the sites of exchange. The O isotope signature of transpired H_2O δ_{w-e}^{18} did not change significantly with temperature (Fig. S3) but the isotopic signature of H_2O at the site of evaporation δ_{w-e}^{18} significantly decreased with temperature (Fig. 4a).

Therefore, the difference between $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ could also be explained by errors in parameterizing the $\delta^{18}\text{O}$ of H_2O at the sites of evaporation (δ_{w-e}^{18}). For example, assuming that b is constant at 29‰ for $\Delta^{13}\text{C}-g_m$ and $\theta = 1$ across

temperatures then the δ_{w-c}^{18} needed to minimize the difference between $\Delta^{18}\text{O-}g_m$ and $\Delta^{13}\text{C-}g_m$ resulted in significantly lower values at leaf temperatures $\geq 30^\circ\text{C}$ (Fig. 4a). Alternatively, assuming that b is constant at 29‰ for $\Delta^{13}\text{C-}g_m$ and δ_{w-c}^{18} is correct then θ can be solved for to minimize differences between $\Delta^{18}\text{O-}g_m$ and $\Delta^{13}\text{C-}g_m$ across leaf temperatures. This caused θ to significantly differ from one at leaf temperature $\geq 30^\circ\text{C}$ ($P_{5,14} < 0.001$) (Fig. 4b).

pH sensitivity of CA activity at chloroplastic $[\text{CO}_2]$

The pH sensitivity of k_{CA} from pH 6.8 to 8.2 measured with the membrane inlet mass spectrometer showed an exponential increase with pH: $k_{CA} = 5 \times 10^{-8} e^{2.1019 \times \text{pH}}$ (Fig. 5). At $> 25^\circ\text{C}$ the CA_{leaf} derived using the k_{CA} at pH 8.0 and C_{c13} , increased with temperature (Fig. 6). This assumes pH remains constant with temperature; however, a shift in pH with temperature would have a significant influence on CA_{leaf} . For example, across all temperatures the modeled CA_{leaf} was significantly lower at pH 7.8 and significantly higher at pH 8.2 compared with pH 8.0 (Fig. 6; dashed vs dotted lines, respectively).

Discussion

Temperature dependency of $\Delta^{13}\text{C-}g_m$ and $\Delta^{18}\text{O-}g_m$

In the current study, net CO_2 assimilation A_{net} , stomatal conductance g_s , and $\Delta^{13}\text{C-}g_m$ in *P. bisulcatum* increased with leaf temperature, similar to that reported by von Caemmerer & Evans (2015) for a large number of C_3 species. The short-term temperature response of $\Delta^{13}\text{C-}g_m$ has been attributed to changes in CO_2 diffusion through the liquid phase (including cell wall, cytoplasm, and chloroplast stroma) and the membrane phase (including the plasma membrane and chloroplast envelopes) (Evans & von Caemmerer, 2013; von Caemmerer & Evans, 2015). However, precise quantification of $\Delta^{13}\text{C-}g_m$ depends on the choice of fractionation factors and underlying photosynthetic model (Eqn 2) (Flexas *et al.*, 2008; Ubierna & Farquhar, 2014). The current study initially assumed that the fractionation factors associated with CO_2 diffusion a_m , Rubisco carboxylation b , respiration ℓ , and photorespiration f were independent of temperature. Previously, Evans & von Caemmerer (2013) validated these assumptions for the temperature response of $\Delta^{13}\text{C-}g_m$ at 2‰ O in tobacco and subsequently measured $\Delta^{13}\text{C-}g_m$ in multiple C_3 species (von Caemmerer & Evans, 2015). As will be discussed shortly, these assumptions are further analyzed to determine if they can reconcile the differences between $\Delta^{13}\text{C-}g_m$ and $\Delta^{18}\text{O-}g_m$ in Fig. 2.

There is less information in the literature on measurements of $\Delta^{18}\text{O-}g_m$, particularly the response of $\Delta^{18}\text{O-}g_m$ to short-term changes in environmental conditions such as temperature. In the *P. bisulcatum* data presented here, at $< 30^\circ\text{C}$ the $\Delta^{18}\text{O-}g_m$ and $\Delta^{13}\text{C-}g_m$ were not significantly different; however, at $\geq 30^\circ\text{C}$ the $\Delta^{13}\text{C-}g_m$ was significantly higher than $\Delta^{18}\text{O-}g_m$. Previous studies have reported species variation in difference between the $\Delta^{18}\text{O-}g_m$ and the $\Delta^{13}\text{C-}g_m$ (Gillon & Yakir, 2000; Barbour *et al.*,

2016; Loucos *et al.*, 2017). For example, Barbour *et al.* (2016) reported a nonsignificant difference between $\Delta^{18}\text{O-}g_m$ and $\Delta^{13}\text{C-}g_m$ in wheat but observed 80% higher $\Delta^{18}\text{O-}g_m$ than $\Delta^{13}\text{C-}g_m$ in tobacco, and $\Delta^{18}\text{O-}g_m$ was more than double $\Delta^{13}\text{C-}g_m$ in cotton at leaf temperatures between 31.1 and 33.8°C. It remains unclear

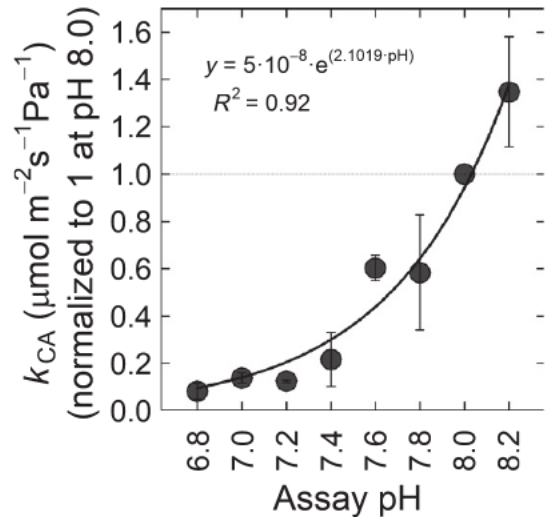


Fig. 5 The pH response of the rate constant for carbonic anhydrase hydration (k_{CA}) for *Panicum bisulcatum* measured by membrane inlet mass spectrometer at 25°C . The k_{CA} values are normalized to the measured value at pH 8.0 (dotted line). Circles are the means of three extractions from three separate plants \pm SD. The solid line is the modeled pH responses using the equation shown in the graph.

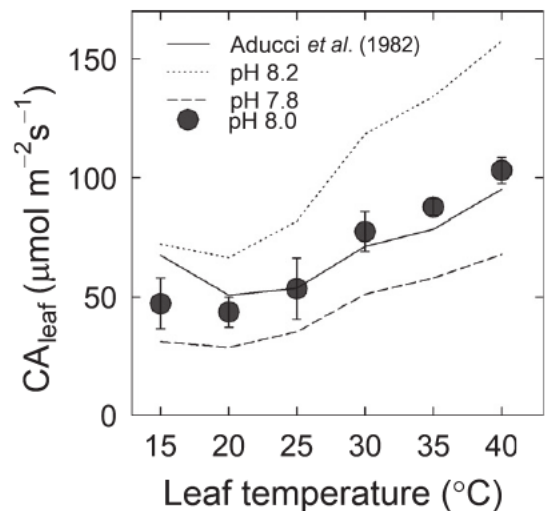


Fig. 6 Temperature response of carbonic anhydrase (CA) activity at chloroplast $[\text{CO}_2]$ in *Panicum bisulcatum*. Chloroplast $[\text{CO}_2]$ was calculated using $\Delta^{13}\text{C}$ method, and the k_{CA} was measured at 25°C and pH 8.0 ($4.5 \mu\text{mol m}^{-2} \text{s}^{-1} \text{Pa}^{-1}$) and modeled temperature response according to Boyd *et al.* (2015) using a modified Arrhenius model, where $E_a = 40.9 \text{ kJ mol}^{-1}$, $\Delta S = 0.21 \text{ kJ mol}^{-1} \text{K}^{-1}$, and $\Delta H = 64.5 \text{ kJ mol}^{-1}$. Circles are the means of four biological replicates \pm SD, $n = 4$. Modeled lines are leaf CA activity CA_{leaf} calculated using the measured pH response of k_{CA} in Fig. 5 ($k_{CA} = 5 \times 10^{-8} e^{2.1019 \times \text{pH}}$) with predicted temperature response of pH (solid line) and constant pH values (dashed and dotted lines).

that measurements of $\Delta^{18}\text{O}-g_m$ and $\Delta^{13}\text{C}-g_m$ can effectively partition g_m into its biochemical and anatomical components.

There are several assumptions needed to estimate $\Delta^{18}\text{O}-g_m$, including that: (1) the $\text{CO}_2\text{-H}_2\text{O}$ exchange occurs at chloroplast surface; (2) there is a full isotopic equilibrium between CO_2 and H_2O at the site of exchange ($\theta = 1$); and (3) the H_2O at the site of exchange δ_{w-cc}^{18} is isotopically similar to the H_2O at the sites of evaporation δ_{w-e}^{18} (Gillon & Yakir, 2000; Barbour *et al.*, 2016). Additionally, it is generally assumed that $\Delta^{18}\text{O}-g_m$ does not incorporate the resistance imposed by the chloroplast membrane and stroma, as is assumed for the $\Delta^{13}\text{C}-g_m$. Specifically, $\Delta^{18}\text{O}-g_m$ provides an estimate of the internal conductance of CO_2 to the site $\text{CO}_2\text{-H}_2\text{O}$ exchange at the chloroplast surface, whereas $\Delta^{13}\text{C}-g_m$ estimates mesophyll conductance from the intercellular airspaces to the site of Rubisco fixation of CO_2 within the chloroplast stroma. Therefore, the $\Delta^{13}\text{C}-g_m$ has been suggested to be *c.* 0.66 of $\Delta^{18}\text{O}-g_m$ (Yakir, 1998) and can be used to separate CO_2 conductance of the chloroplast g_{ch} (i.e. chloroplast envelope and stroma) and wall g_w (i.e. cell wall, plasma membrane. and cytosol) (Gillon & Yakir, 2000). However, in addition to these potentially inherent differences between these estimates of g_m , the differences between $\Delta^{18}\text{O}-g_m$ and $\Delta^{13}\text{C}-g_m$ can be significantly influenced by the input parameters used in their calculation.

Assuming the calculations of $\Delta^{18}\text{O}-g_m$ and $\Delta^{13}\text{C}-g_m$ for *P. bisulcatum* data presented here were correctly parameterized at 25°C suggests one or more of the following: (1) the resistance to CO_2 diffusion lies entirely within the cell wall and plasma membrane; (2) the assumption for the $\text{CO}_2\text{-H}_2\text{O}$ exchange at the chloroplast surface may be incorrect; and (3) there might be some flaws in the assumptions for the estimation of g_m in both methods (i.e. the $\Delta^{18}\text{O}-g_m$ and the $\Delta^{13}\text{C}-g_m$). However, it is unlikely that the cell wall and plasma membrane provide the only resistance to CO_2 movement into the chloroplast, because the double membrane surrounding the chloroplast must impose some resistance to CO_2 diffusion (Uehlein *et al.*, 2008).

It should be noted that although $\Delta^{18}\text{O}-g_m$ and $\Delta^{13}\text{C}-g_m$ differed in sensitivity to temperature, both estimates increased with increasing temperature. The thermal sensitivity of CO_2 conductance in the liquid phase is thought to be limited, whereas CO_2 conductance through membranes increases exponentially with temperature (Evans & von Caemmerer, 2013). The observation of a thermal response for $\Delta^{18}\text{O}-g_m$ suggests that the sites of $\text{CO}_2\text{-H}_2\text{O}$ equilibration must lie interior to at least one membrane. However, as parameterized, the $\Delta^{18}\text{O}-g_m$ measured in the current study was not as temperature sensitive as $\Delta^{13}\text{C}-g_m$. As will be discussed shortly, the difference in the temperature response of $\Delta^{18}\text{O}-g_m$ and $\Delta^{13}\text{C}-g_m$ may be due to errors in the assumption used in the calculations.

$\Delta^{18}\text{O}-g_m$

To our knowledge, there are no reports investigating the temperature response of $\Delta^{18}\text{O}-g_m$ in C_3 species; hence, uncertainty in the assumptions associated with $\Delta^{18}\text{O}-g_m$ with changing temperature remained unexplored. Therefore, we tested these assumptions with the caveat that $\Delta^{18}\text{O}-g_m$ measures g_m to the

chloroplast surface and that $\Delta^{13}\text{C}-g_m$ estimates g_m to the site of carboxylation such that the inherent difference between $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ is primarily determined by the conductance across the chloroplast envelope and can be accounted for as $\Delta^{13}\text{C}-g_m \approx 0.66 \times \Delta^{18}\text{O}-g_m$ across temperatures (Yakir, 1998). Under this scenario, the modeled θ was significantly < 1 at $\geq 30^\circ\text{C}$. Alternatively, assuming that θ was constant across temperatures, the modeled δ_{w-e}^{18} required to minimize the difference between $\Delta^{13}\text{C}-g_m$ and $\Delta^{18}\text{O}-g_m$ was significantly higher at $\geq 30^\circ\text{C}$ than the calculated δ_{w-e}^{18} .

The assumption that the isotopic composition between the H_2O at the site of exchange δ_{w-cc}^{18} is the same as that at the sites of evaporation δ_{w-e}^{18} depends on the spatial separation of these two locations within the leaf and the potential spatial variation of the isotopic signature of H_2O within the leaf. Alternatively, the assumption that there is a full isotopic equilibrium between CO_2 and H_2O at the site of exchange (e.g. $\theta = 1$) is primarily determined by the activity of CA. Moreover, the temperature effect on CA activity and differences between δ_{w-cc}^{18} and δ_{w-e}^{18} must also be taken into account.

Leaf temperature may affect CA activity due to temperature-mediated changes in CA catalytic properties and potentially shifts in the cytosolic/chloroplastic pH. Additionally, deactivation of CA in *S. viridis* at temperatures $> 25^\circ\text{C}$ was reported by Boyd *et al.* (2015), suggesting that the CA activity may limit θ at higher temperatures, particularly if chloroplast $[\text{CO}_2]$ also decreases. However, the increase of $\Delta^{13}\text{C}-g_m$ with temperature suggests that chloroplast $[\text{CO}_2]$ also increases with temperature, potentially offsetting any deactivation of CA. Furthermore, at 25°C we observed an exponential increase in CA hydration rate k_{CA} from pH 6.8 to 8.2 (Fig. 5), similar to that previously published by Berg *et al.* (2015). If pH changed with temperature, as previously reported by Aducci *et al.* (1982), who saw a 0.5 unit decrease in the cytosolic pH of maize root tip tissue with increasing temperature from 4 to 28°C, this would significantly influence not only CA activity but also potentially other reactions. Unfortunately, to our knowledge, there are no reports addressing temperature dependency of cytosolic or chloroplast pH in a photosynthetically active leaf, and to measure *in vivo* pH is technically beyond the scope of the present study.

However, modeling CA activity in response to temperature at both pH 8.2 and pH 7.8 demonstrates that a relatively small change in pH could have a significant influence on CA_{leaf} . Therefore, a potential decrease in cytosolic/chloroplastic pH with temperature may lead to a decrease in θ due to reduced CA activity. It is worth noting that CA_{leaf} in *P. bisulcatum* is low relative to other C_3 species (Gillon & Yakir, 2000). The low CA activity in *P. bisulcatum* may lower θ , particularly at high leaf temperatures, and this response might not be as pronounced in species with higher CA_{leaf} . Therefore, future studies on the temperature response of CA activity in C_3 species with diverse levels and/or anti-CA lines of tobacco, Arabidopsis, and rice are needed.

It is also possible that the isotopic signature of H_2O at the sites of evaporation δ_{w-e}^{18} as calculated with the Craig–Gordon model (Eqn 7) does not accurately represent the signature of H_2O at the site of exchange δ_{w-cc}^{18} . In C_3 plants, it has been proposed that the

site of O exchange between leaf H₂O and CO₂ is primarily located at the chloroplast surface (Yakir, 1998; Fabre *et al.*, 2007). The majority of chloroplast in the mesophyll cells in C₃ plants is appressed to the cell walls adjacent to the intercellular airspace, so it is generally assumed that δ_{w-e}^{18} is a good approximation of δ_{w-ce}^{18} (Gillon & Yakir, 2000; Barbour *et al.*, 2016). This assumes that the site of evaporation occurs near the cell walls next to the intercellular airspace in close proximity to the chloroplast. However, the site of H₂O evaporation within the leaf is not specifically known and may occur relatively far from the mesophyll cells adjacent to guard cells (Sack & Holbrook, 2006; Buckley *et al.*, 2017). Furthermore, it has been long recognized that bulk leaf H₂O is more depleted than δ_{w-e}^{18} due to the combined contribution of source H₂O and the H₂O at the sites of evaporation. Several models have been developed (e.g. the two-pool and the Pécelet models) to describe how the isotopic composition of bulk leaf H₂O is influenced by source H₂O and the H₂O at the sites of evaporation (Farquhar & Lloyd, 1993; Gillon & Yakir, 2000; Barbour & Farquhar, 2003; Tomás *et al.*, 2013; Barbour *et al.*, 2016; Holloway-Phillips *et al.*, 2016). However, it remains unclear what type of isotopic gradient might occur within a transpiring leaf, making it difficult to precisely parameterize the δ_{w-ce}^{18} from measurements of δ_{w-e}^{18} , particularly if the site of exchange is relatively distant from the sites of evaporation. This uncertainty is compounded by the fact that the location of O exchange between CO₂ and leaf H₂O is unknown and may differ as rates of transpiration and leaf temperature change. In fact, Barbour *et al.* (2016) demonstrated that estimates of $\Delta^{18}\text{O}-g_m$ are significantly sensitive to changes in the Pécelet effect. Taken together, it remains unclear how to effectively parameterize θ and the discrepancies between δ_{w-ce}^{18} and δ_{w-e}^{18} when estimating $\Delta^{18}\text{O}-g_m$. However, it appears clear that the assumptions used at 25°C likely do not hold true for higher temperatures.

Conclusion

We have estimated temperature responses of $\Delta^{13}\text{C}-g_m$ and the $\Delta^{18}\text{O}-g_m$ using coupled leaf gas exchange and isoflux measurements of CO₂ and transpired H₂O in C₃ *P. bisulcatum*. Our observations are unable to partition g_m into its components (i.e. g_w and g_{ch}) and their temperature dependency because of uncertainties in the temperature response of several input parameters. However, the data presented here suggest that the highest uncertainties are associated with the assumptions made in calculating $\Delta^{18}\text{O}-g_m$ (e.g. θ and $\delta_{w-e}^{18} = \delta_{w-ce}^{18}$). Future work to obtain precise information on the temperature dependency of cytosolic and chloroplastic pH could better enable partitioning of g_m into its components and their responses to environmental changes using combined measurements of $\Delta^{13}\text{C}-g_m$ and the $\Delta^{18}\text{O}-g_m$.

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Author contributions

BVS and ABC designed the experiment. BVS performed the measurements. BVS and ABC analyzed the data and wrote the manuscript.

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Supporting Information

Additional Supporting Information may be found online in the Supporting Information section at the end of the article:

Fig. S1 Sensitivity of Picarro (L2130-i) for $\delta^{18}\text{O}$ measurements to the water vapor concentration.

Fig. S2 Temperature response of transpiration rates.

Fig. S3 Corrections of $\delta^{18}\text{O}$ of transpired water ($\delta_{w-E}^{18}\text{O}$) by accounting $\delta^{18}\text{O}$ offset of Picarro (L2130-i).

Methods S1 Calculations for CO₂ mesophyll conductance from $\Delta^{18}\text{O}$ ($\Delta^{18}\text{O}-g_m$).

Table S1 Parameters and units or values for calculation of g_m with the $\Delta^{13}\text{C}-g_m$ method.

Table S2 Parameters and units or values for calculation of g_m with the $\Delta^{18}\text{O}-g_m$ method.

Table S3 Summary of temperature response for gas exchange, isoflux and mesophyll conductance (calculated under standard assumptions) for the youngest fully expanded leaf of *Panicum bisulcatum*.

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